

The Effects of Various Water Table Depths on CO₂ Emission at Oil Palm Plantation on West Aceh Peat

Etik Puji Handayani¹, Meine van Noordwijk², Kamarudin Idris³
Supiandi Sabiham³ and Sri Djuniwati³

Received 2 September 2009 / accepted 19 June 2010

ABSTRACT

The Effects of Various Water Table Depths on CO₂ Emission at Oil Palm Plantation on West Aceh Peat (EP Handayani, M van Noordwijk, K Idris, S Sabiham and S Djuniwati): Changes in the depth of water table influenced carbon cycling in peatlands, and affected the sources and sinks of carbon dioxide. The effects of depth of water tables in oil palm plantations on the emission of CO₂ were studied. CO₂ emissions of peatland were measured in Meulaboh, West Aceh using cylindrical chambers and air samples from the chambers were analyzed by gas chromatography. Five-point transects perpendicular to drainage canals provided variation in the depth of water tables for the samples. Data from oil palm fields were compared to data from an adjacent swamp forest. The data confirmed that the increasing depth of water table was accompanied by the increasing in microbial activity that was measured by CO₂ emission. The CO₂ emissions from chambers with additional root zones were higher than from bulk soil chambers between one to four times.

Keywords: CO₂ emissions, depth of water table, oil palm, peat land

INTRODUCTION

Peatlands worldwide play a vital role in biosphere biogeochemical processes. Peatland constitutes an important biome because of its high soil C content. Turenen *et al.* (2002) estimated that peatlands store 220–460 pg of carbon and hence can significantly influence atmospheric CO₂ concentrations (Hilbert *et al.* 2000).

The total area of peat lands in Indonesia is about 20 million ha (Rieley 1996) and the average oil palm yield on peat land can reach 23 Mg ha⁻¹y⁻¹ fresh fruit bunches (Winarna, 2007). Therefore, peat lands have considerable potential for development of oil palm agrobussines in Indonesia. However, peat lands contain one-third of global soil carbon and total stocks represent 70 years of current annual global emission from fossil fuel burning. This carbon store is now

being released to the earth's atmosphere through forest fire and respiration, both increased by drainage.

Like many ecosystems, peat land accumulates C under undrained (undisturbed) conditions and emits C (CO₂ by oxidization) under drained (disturbed) conditions (Nykänen *et al.* 1995, Laiho *et al.* 1996, Silvola *et al.* 1996). Peat soil subsidence after drainage is often seen as a rough measure for CO₂ emissions, though little agreement exists on the fraction of the subsidence that can be attributed to oxidization.

Carbon dioxide gas is part of the greenhouse effect on global warming. CO₂ gas emissions vary with stages of plant growth, depending on management practices for soil and plants (such as drainage and fertilization) and characteristics of peat land, including water level, and the thickness and maturity of the peat deposits.

¹STIPER Dharma Wacana Metro, Lampung, Indonesia, email etik_ph@yahoo.com;

²World Agroforestry Centre, South East Asia Office, Bogor, Indonesia;

³Bogor Agricultural University, Bogor, Indonesia.

This research was aimed to be assessing the influence of depth of water tables on CO₂ emissions on the areas with roots and without roots.

MATERIALS AND METHODS

Study Site

This research was conducted in the peat domes of Aceh Barat District, Nanggroe Aceh Darussalam (NAD) Province of Indonesia in 2008. The emitted CO₂ gas was captured using closed chambers from the rooted (R) and non-rooted (NR) chambers or zone. The observation was conducted at smallholder oil palm plantations with 10 and 5 year old plant stands. For each of the age groups, three to five pairs of observation points were made one transect which have the same depth of peat, each point with R and NR treatment.

Analysis of CO₂

The closed chambers of 30 cm diameter and 30 cm tall were made of PVC tubes. The bottom brim of the chamber was sharpened to minimize soil compaction during their insertion into the ground. For each chamber designed for rooted zone emission, a hole of 5 cm diameter at a point 20 cm from its top was made for channelling three pieces of oil palm roots in such a way that the roots can still grow and develop inside the chambers. These chambers were installed at a distance of 2.5 m from the trunk of oil palms aged 10 years, and 1 m from the trunk of oil palm aged 5 years according to the distribution of the plant roots. The paired R and NR chambers were

mounted at a distance of 1 m from each other. Each chamber was equipped with a septum to place the needle puncture. A small (6 cm battery powered) fan was installed inside the chambers to stir the gas. A thermometer was also installed for each chamber to measure the temperature during the gas sampling.

Time of Observation

Gas samples were taken by using syringes of 5 and 10 ml capacity, with a sampling frequency of 0, 5, 10, 15, 25 and 35 minutes after closing of the chambers. All sampling was conducted during the morning hours from 07:00 to 10:00 a.m. for best consistency. Samples were analyzed within 24 hours after sampling using a portable gas chromatography instrument. Measurements were conducted from October to November 2008 to represent early rainy season. For each gas emission measurement point, the depth of the water table was also measured from a hole as deep as 1 m made with a soil auger at the midpoint between the R and NR.

RESULTS AND DISCUSSION

The rate of CO₂ emissions in the peat soils was in the range of 10 - 40 Mg ha⁻¹ yr⁻¹. The CO₂ emissions from each transect is shown in Figures 1 - 7.

Carbon dioxide emission in rhizosphere zones (the chamber with roots) were higher than non rhizosphere zones (the chamber without roots). However, the effect of water table depth on CO₂ emissions was rather inconsistent. Other researchers demonstrated a positive correlation between CO₂

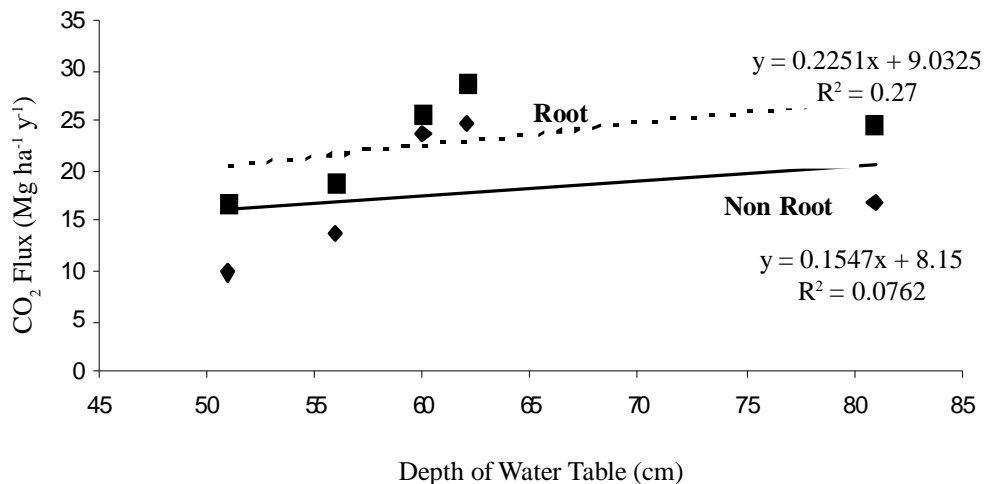


Figure 1. CO₂ emission under a smallholder oil palm plantation in Suak Puntong, Transect 1.

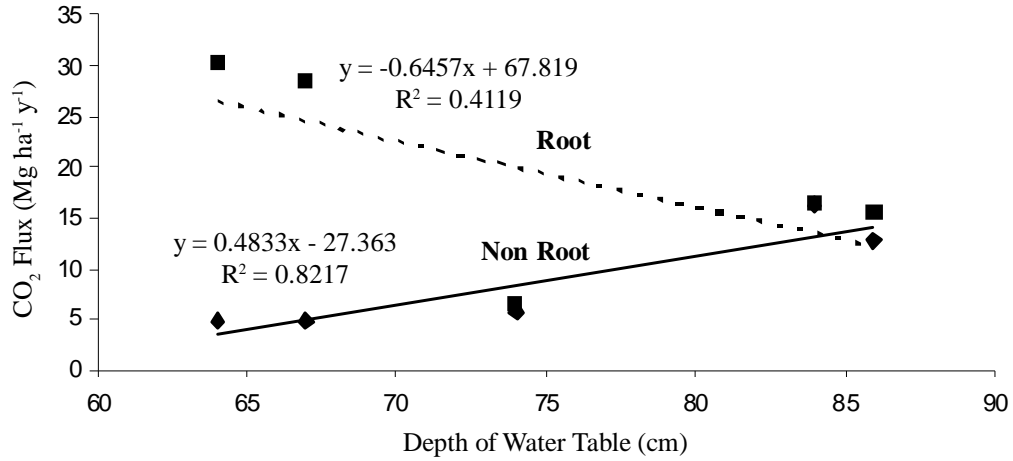


Figure 2. CO₂ emission under a smallholder oil palm plantation in Suak Puntong, Transect 2.

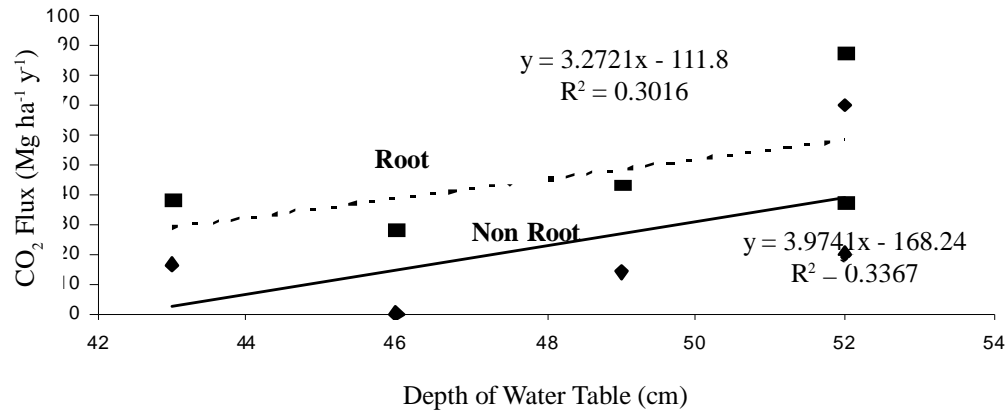


Figure 3. CO₂ emission under a smallholder oil palm plantation in Suak Raya, Transect 3.

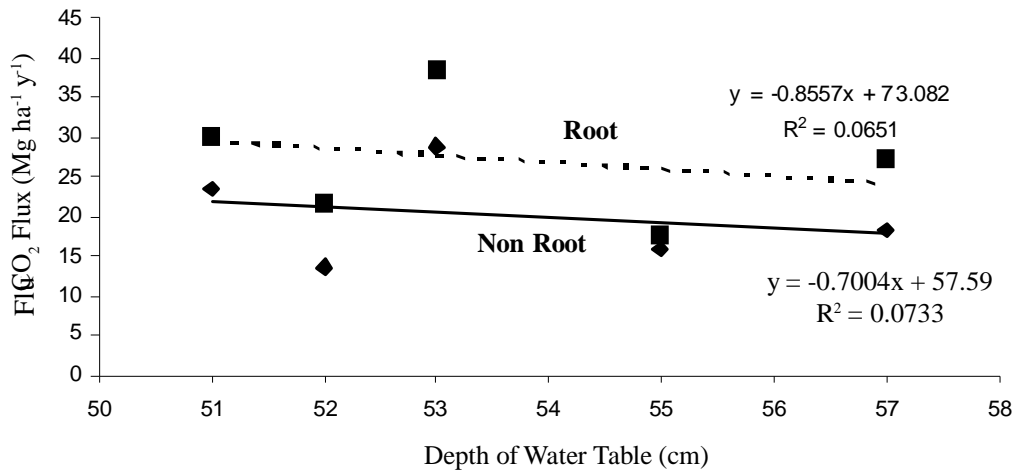


Figure 4. CO₂ emission under a smallholder oil palm plantation in Suak Raya, Transect 4.

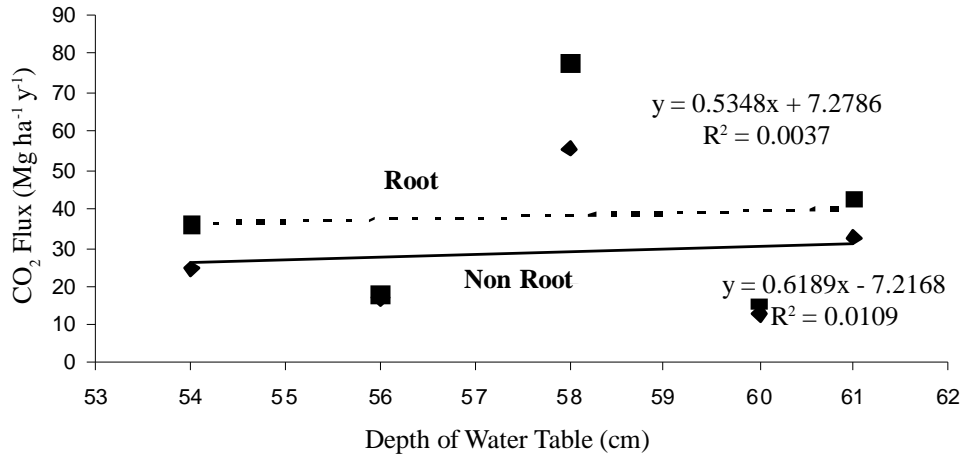


Figure 5. CO₂ emission under a smallholder oil palm plantation in Suak Raya, Transect 5.

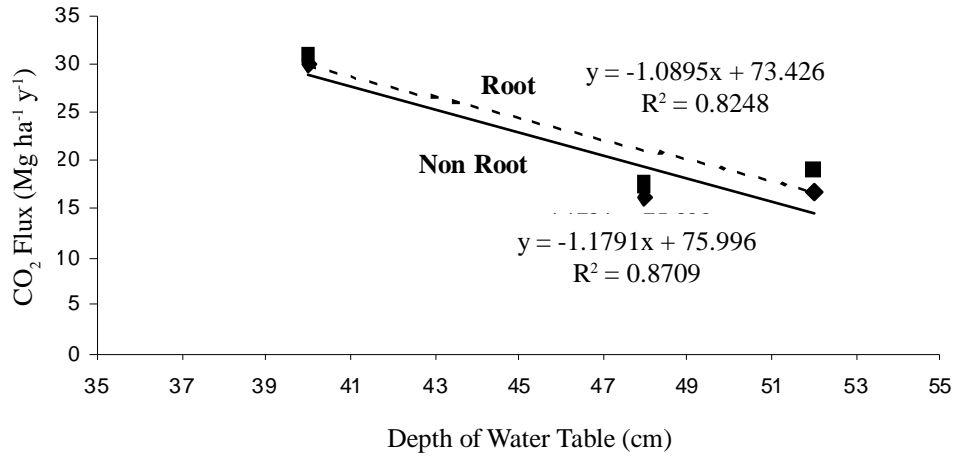


Figure 6. CO₂ emission under a smallholder oil palm plantation in Suak Raya, Transect 6.

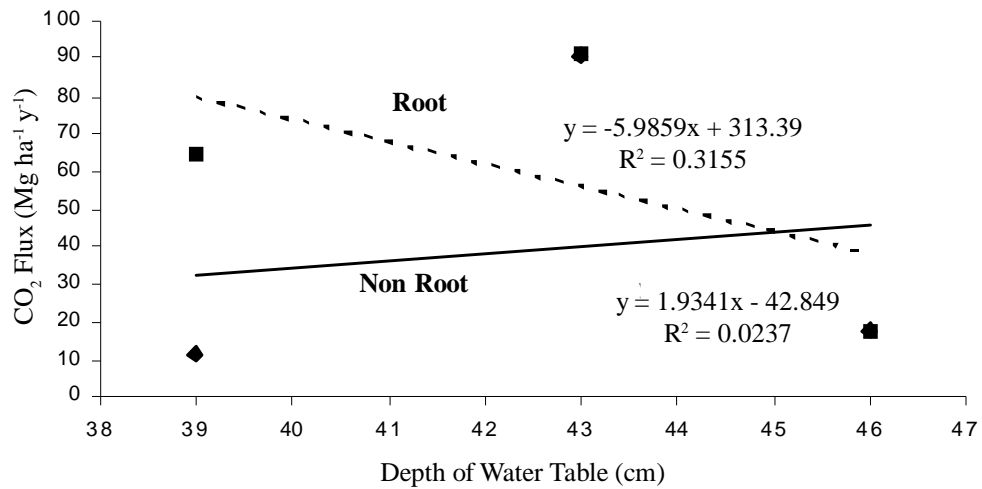


Figure 7. CO₂ emission under a smallholder oil palm plantation in Suak Raya, Transect 7.

emissions and water table depth (Hooijer *et al.* 2006; Jauhiainen *et al.* 2001). The high variation in our data may have contributed to these inconsistencies. At transect 1, 2, 3, 5 and 7 the phenomenon of CO₂ emission increased with increasing depth of water table, while at transect 4 and 6 would be decreased. CO₂ emission decreased with increasing depth of water table at transect 4 and 6 due to the peat thickness only 142-170 cm than the other transect. Yulianti (2009) reported that the reserved of C peat had a positive correlation to thickness of peat ($r = 0.93$).

Results of measurement of depth of water table in each transect indicate that in the same transect, the point of measurement which was close to the drainage channel had the deeper water table, and progressively far from the drainage channel, hence depth of water table decreased. All of the transect except transect SR-4 and SR-6 showed the same correlation that was progressively far from the drainage channel, the water tables were shallow and CO₂ emissions decreased. It was because the nearer point of measurement with the drainage channel, deeper water table caused advance decomposition. Condition aeration would be faster to improve the Oxygen availability in peat substances and could be accelerated the process of C-Organic mineralization. As a results, peat substances produced CO₂, so that CO₂ emissions were higher in the point of measurement near from the drainage canals. Thereby, it could be said that the depth of the water table represents one of factor influencing CO₂ emission. Other researchers reported that decomposition was very influenced by depth of irrigation and water fluctuation irrigation (Belyea and Clymo 2001), so that the content irrigated to influence the emission CO₂ from underground (Smith *et al.* 2003; Liu *et al.* 2008). Silva *et al.* (2008) reported that emission of CO₂ 1.2 times bigger at soil which incubation with water holding capacities 40%, 60% and 80% compared to soil with water holding capacities 100%. According to Jia *et al.* (2006), too high water content would pursue the diffusion of CO₂ and microbial activity. However Stark *et al.* (2004) also stated that lowering water content would degrade the microbial activity.

In this research, the value of non root in closed chamber emission, with mean ranging from 18 to 24 Mg CO₂⁻¹ha⁻¹yr⁻¹ were less half than the predicted values of Hooijer *et al.* (2006) of around 54 Mg CO₂⁻¹ ha⁻¹ y⁻¹ under the depth drainage (in this case, the water table) of around 60 cm and Melling *et al.* (2005) of around 50 Mg CO₂⁻¹ ha⁻¹ yr⁻¹ for oil palms in

Sarawak Malaysia and Jauhiainen *et al.* (2001) for stabilized agricultural land. This was because of the limited measurements conducted under the current experiment. Nevertheless the effect of rhizosphere was consistent under the current experiment. Hirano *et al.* (2007) expressed that result of measurement of CO₂ emissions from peat tropic was very high variation depends on time and measurement place, when farms start in conversion, differences microclimate such as temperature of under ground and air temperature, status of nutrient.

Carbon dioxide emissions from rhizosphere zones were 1 up to 4 times bigger than non rhizosphere zones. It was because of the quality of root media capable of changing the nature of physical, chemical and soil biology in an around roots (Darrah 1991; Gregory and Hinsinger 1999). Rhizosphere had the conducive environment expanding a lot of organisms (Bowen and Rovina 1973; Peterson 2003) hence a lot of processes in the rhizosphere directly and also indirectly increased capacities of soil function for the growth of crops and as an environmental buffer (Gregory and Hinsinger 1999). The total of microbial activity as an effect of height of concentration nutrias, C-Labile, and exudates grow on around root area (Kuzyakov *et al.* 2000; Subke *et al.* 2004; Hamer and Marschner 2005). Thereby, produce the CO₂ representing resultants from respiration of microorganism and the value of root respiration on the rhizosphere that were higher than that on the non rhizosphere.

CONCLUSIONS

CO₂ emission increased with increasing the depth of the water table. However, the reverse pattern was also found, and another pattern where CO₂ emission was independent the depth of the water table. It would cause a strong over estimate of peat soil contribution to CO₂ emission when the measurement was conducted in rooted area. Thus when zoning of the measurement was not possible, measurement should be conducted on a relatively root free areas. The CO₂ emissions from chambers with additional root zones were one to four times higher than from bulk soil chambers.

REFERENCES

- Belyea LR and RS Clymo. 2001. Feedback control of the rate of peat formation. *Proc Royal Soc* 268: 1315-1321.

- Bowen GD and AD Rovira. 1973. Are modelling approaches useful in rhizosphere biology. *Bull Ecol Res Com* 17: 443-450.
- Darrah PR. 1991. A models of the rhizosphere. I: Microbial population dynamics around a root releasing soluble and insoluble carbon. *Plant Soil* 133: 187-199
- Gregory PJ and P Hinsinger. 1999. New approaches to studying chemical and physical changes in the rhizosphere: an overview. *Plant Soil* 211: 1-9.
- Hamer U and B Marschner. 2005. Priming effects in soils after combined and repeated substrate additions. *Geoderma* 128: 38-51.
- Hilbert DW, NT Roulet and T Moore. 2000. Modelling and analysis opeatlands as dynamical systems. *J Ecol* 88: 230-242.
- Hirano T, H Segah, T Harada, S Limin, T June, R Hirata and M Osaki. 2007. Carbon dioxide balance of a tropical peat swamp forest in Kalimantan, Indonesia. *Global Change Biol* 13: 1-14.
- Hooijer A, M Silvius, H Wosten and S Page. 2006. PEAT-CO₂, Assessment of CO₂ emissions from drained peatlands in SE Asia. Wageningen: Delft Hydraulics report Q3943.
- Jauhiainen J, J Heikkinen, PJ Martikainen and H Vasander. 2001. CO₂ and CH₄ fluxes in pristine peat swamp forest and peatland converted to agriculture in central Kalimantan, Indonesia. *Int Peat J* 11: 43-49.
- Jia B, G Zhou and Y Wang. 2006. Effects of temperature and soil water-content on soil respiration of grazed an ungrazed leymus chinensis steppes, Inner Mongolia. *J Arid Environ* 67: 60-76.
- Kuzyakov Y, JK Friedel and K Stahr. 2000. Review of mechanisms and quantification of priming effects. *Soil Biol Biochem* 32: 1485-1498.
- Laiho R, J Laine and H Vasander. 1996. Northern peatlands in global climatic change. In: Proceedings of the International Workshop, Hytiala, Finland, 8-12 October 1995.
- Liu LC, Y Fan, G Wu and YM Wei. 2008. Using LMDI method to analyze the change of China's industrial CO₂ emissions from final fuel use: an empirical analysis. *Energy Policy* 35 (11): 5892-5900.
- Melling L, R Hatano and KJ Goh. 2005. Global Warming Potential from peatland of Sarawak, Malaysia. *Phyton (Austria) Special issue* 45: 275-284.
- Nykänen H, J Alm, K Lang, T Silvola and PJ Martikainen. 1995. Emissions of CH₄, N₂O and CO₂ from a virgin fen drained for grassland in Finland. *J Biogeography* 22: 351-357.
- Peterson E. 2003. Importance of rhizodeposition in the coupling of plant and microbial productivity. *European J Soil Sci* 54: 741-750.
- Rieley JO, AA Ahmad-Shah and MA Brady. 1996. The extent and nature of tropical peat swamps. In: E Maltby, CP Immirzi and RJ Safford (eds) *Tropical Lowland Peatlands of Southeast Asia*. IUCN, Gland, pp.17-53.
- Silva CC, ML Guido, JM Cebbalos, R Marsch and L Dendooven. 2008. Production of carbon dioxide and nitrous oxide in alkaline saline soil of Texcoco at different water content amended with urea: A laboratory study. *Soil Biol Biochem* 40: 1813-1822.
- Silvola J, J Alm, U Ahlholm, H Nykänen and PJ Martikainen. 1996. CO₂ fluxes from peat in boreal mires under varying temperature and moisture conditions. *J Ecol* 84: 219-228.
- Smith KA, T Ball, F Conen, KE Dobbie, J Massheder and A Rey. 2003. Exchange of greenhouse gases between soil and atmosphere: Interaction of soil physical factors and biological processes. *European J Soil Sci* 54: 779-791.
- Strack M, JM Waddington and ES Tuittila. 2004. Effect of water table drawdown on northern peatland methane dynamics: Implications for climate change. *Global Biogeochem Cycl* 18: 4003-40010.
- Subke JA, V Hahn, G Battipaglia, S Linder, N Buchmann and MF Cotrufo. 2004. Feedback interactions between needle litter decomposition and rhizosphere activity. *Oecologia* 139: 551-559.
- Turenen J, E Tomppo, K Tolonen and A Reinikainen. 2002. Estimating carbon accumulation rates of undrained mires in Finland—applications to boreal and subarctic regions. *Holocene* 12: 69-80.
- Winarna. 2007. Lahan Gambut Saprik paling Potensial Untuk Kebun Sawit. <http://www.kapanlagi.com/h/old/0000179066.html> [Accessed on 25-8-2007], (in Indonesian).
- Yulianti N. 2009. Cadangan karbon lahan gambut dari agroekosistem kelapa sawit PTPN IV Ajamu, Kabupaten Labuhan Batu, Sumatera Utara. [Tesis]. Bogor: Program Pascasarjana, Institut Pertanian Bogor (in Indonesian).